

Perspective: The Great AMOC Shutdown

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1 A Perspective

Much discussion of the influence of the changing ocean on climate past and present, invokes the idea of a “shutdown” of the Atlantic meridional overturning circulation (Atlantic MOC, or AMOC. See e.g., Cronin, 2010, Stammer et al., 2019 and numerous references to be found there). The specific definition of the AMOC is often omitted and thus the meaning of a “shutdown” mantra remains obscure. The purpose of this note will be regarded by some as simplistic: to point out that a shutdown-AMOC in most implied usage nonetheless requires: (1) a remaining massive transport of fluid and its properties and (2) curiously, is most simply interpreted as an ocean circulation conforming to the classical laminar gyre circulation picture of a Sverdrup-interior plus western boundary current.

A deprecated device in ancient Greek drama is labelled as a *deus ex machina* or ‘god from a box’ (Oxford English Dictionary), a theatrical device in which a god emerges from a container at the end of the a play resolving all plot and human difficulties. In climate and paleoclimate, the idea of a shutdown in the so-called Atlantic Meridional Overturning Circulation (AMOC) is meant to explain many observations, past, present, and future, and to describe the impact on the wider observational system. An arbitrary, but recent, example is Stocker (2021).

To a great extent, the emphasis on the AMOC stems from a cartoon picture of the ocean circulation by the late Wallace Broecker—what he called the "Great Ocean Conveyor", and his invocation of a zoomorphic attribute "the climate is an angry beast..."; recently reproduced as part of a New York Times story. Much of the scientific literature purporting to describe the oceanic past, present, and future, simply invokes the idea of “the” AMOC without ever bothering to define it. A useful working definition is essential, even if it need not be a unique

26 one. A general definition, applying to any ocean, zonally bounded or otherwise, is the meridional
27 overturning circulation or the sum of the mass flux from a western to an eastern boundary of
28 the ocean, to some specified depth (usually not the bottom) at some specified latitude of the
29 northward and southward going flows.¹ Thus the meridional overturning circulation (MOC) is
30 the net flow going north or south above the integration depth (which is potentially a complicated
31 curve). If an ocean is bounded e.g., at the north, the integral (sum) from top to bottom would
32 have to vanish, and thus the MOC is normally defined in terms of some finite depth, not the
33 bottom, perhaps varying with longitude, and definitely varying with latitude. If unbounded
34 zonally, as in some latitudes of the Southern Ocean, the integral can be over 360° zonally, or
35 over any finite longitudinal interval. In a useful definition, a temporal averaging time is specified
36 as well.

37 In the Atlantic Ocean, various definitions of the AMOC exist, generally all referring to the
38 net northward movement of mass above depths of order 1000m. The major, permanent feature,
39 of the North Atlantic Ocean is the powerful *poleward* flow on the western side, known as the
40 Gulf Stream—a western boundary current (WBC) which is a fundamental phenomenon of all
41 ocean basins bounded on the west. The North Atlantic is nearly closed at its northernmost
42 reaches (a weak mass *input* exists there from the Arctic Sea) and the far larger amount of water
43 headed northwards in the Gulf Stream at e.g., 30°N and definable with different numbers at
44 other latitudes, must return southward in the ocean further east.

45 To the extent that a meridional mass transport imbalance exists above the integration depth,
46 inference is made of a convective vertical exchange motion at high latitudes, with the balancing
47 fluid returning at great depths. Recall however, that Scott and Marotzke (2002) pointed out
48 that the conversion in ocean circulation models between the nominal upper layer and the oceanic
49 abyss occurred not as a mass advection, but instead as an enhanced vertical mixing rate localized
50 at the horizontal boundary.

51 A large heat transport can exist by virtue of the balancing return flow having lower tem-
52 peratures than fluid carried northward by the WBC. The “shutdown” of this system is best
53 understood in terms of the two-different eras of thought concerning the ocean circulation gen-
54 erally.

55 *Classical and Modern Periods*

56 Understanding and study of physical oceanography over the years led to views that are use-
57 fully divided into two periods: a classical one prior to about 1975, and a modern one thereafter.²
58 In the classical period, driven by the available observational capability, the ocean was widely

¹The equivalent volume flux in the Boussinesq approximation.

²A “post-modern” period, may some day be defined—but not yet.

59 perceived as a slow-moving, large-scale, laminar, near-steady state, changing at depth only over
60 thousands of years. The classical period gave rise to a remarkable analytical literature of ocean
61 fluid phenomena including Ekman layers, Sverdrup balance, western intensification, Stommel-
62 Arons flows, abyssal recipes, Rossby waves, and numerous other laminar flow regimes. Wunsch
63 and Ferrari (2018) provided a historical sketch and the overall classical picture can be found in
64 any of several readable textbooks.

65 In this classical period, a central paradigm explained the existence of the intense time-mean
66 flows on the western sides of oceans as a consequence of the torque of the large-scale windfield
67 acting on the ocean. The textbook result is a picture (1) of the oceanic “gyres” with strong
68 poleward motions in currents such as the Gulf Stream in the North Atlantic, the Kuroshio in the
69 North Pacific, and the Agulhas in the Indian Ocean, and the Brazil Current, and East Australia
70 Current in the Southern hemisphere oceans. Understanding in this classical picture, dating back
71 to about 1950, is that the massive poleward flow is returned further east in an upper ocean layer
72 usually labelled as “Sverdrup balance”. Any motions at depth (below, very roughly, 1000m) are
73 regarded as negligible.

74 Thus for example, in the North Atlantic at about 30°N, the resulting major northward
75 transport in the Gulf Stream is about $35 \times 10^6 \text{m}^3/\text{s}$ (35 Sverdrups, Sv) $\approx 35 \times 10^3 \text{kg}/\text{s}$. The
76 Gulf Stream exists (as shown by Stommel, 1948 and numerous later authors) as a consequence
77 of the presence of an oceanic western boundary, and a large-scale wind system consisting of the
78 trade winds (easterlies)—and the westerlies to their north. *Barring a major planetary collision,*
79 *no circumstances exist, for tens of millions of years in the past or future, in which this feature*
80 *could or will, have vanished.* A major change in the wind system could strengthen or weaken
81 it, but no known phenomenon is capable of completely eliminating the westerlies and easterlies
82 which are fundamental phenomena of angular momentum conservation in a heated, rotating,
83 atmosphere on a sphere.

84 In a well-known paper from the end of the classical period, Leetmaa et al. (1977) attempted
85 to show that the North Atlantic Ocean was in Sverdrup balance at 16°N, 24°N, 32°N . They
86 assumed, as was then conventional wisdom, that a level-of-no-(horizontal) motion existed near
87 1500m, giving rise to a 2-layer ocean in which the lower layer flow was effectively negligible. Their
88 Gulf Stream carried the requisite mass flux of order 30 Sv northward, but varying considerably
89 with latitude, and with a nearly opposite southward return flow, in the open ocean east of the
90 Gulf Stream, governed by the wind-stress curl, existing in the upper layer.

91 *This picture is one of an ocean in which the AMOC is shutdown:* The zonal integral in
92 the upper ocean vanishes—nothing is returned at depth—but a massive oceanic flow exists,
93 nonetheless (uncertainties are large however). The return flow carries (in their picture) the same

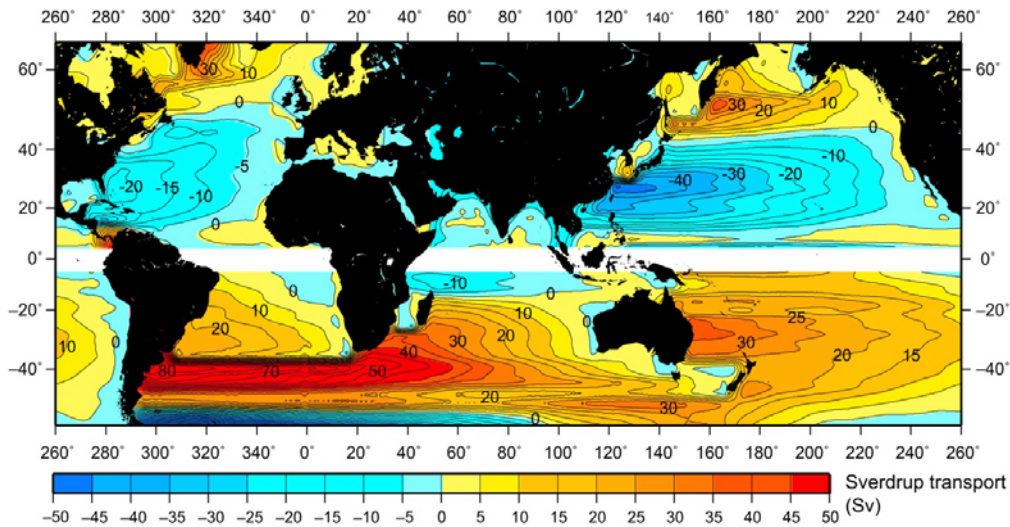


Figure 1: Talley et al. (2011) estimate from a particular estimate of the time-average wind field, of the Sverdrup transport flow. The Gulf Stream and other near-surface western boundary currents (WBCs) are largely fed from the zonal component of the interior Sverdrup flow. But note that sources of fluid in the WBC, can lie also far-distant in the opposite hemisphere. For example, note that the Gulf Stream is fed in part from the South Atlantic (e.g., Schmitz and Richardson, 1991). Equatorial band calculation is omitted for technical reasons.

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94 volume flux, but a different temperature (enthalpy), salt, carbon, etc. content than does the
 95 northward-going Gulf Stream. A net northward North Atlantic flux of heat thus exists—*despite*
 96 *the shutdown*—although it is smaller than existing best estimates (Wunsch and Roemmich,
 97 1985). In a theoretical ocean, the Gulf Stream return flow need not be in Sverdrup balance,
 98 but vorticity conservation of the ocean circulation must still exist. No other simple dynamical
 99 balance is known. AMOC shutdown does not, and cannot, involve a cessation of the Gulf Stream
 100 flow, barring a radical change in the most fundamental Earth physics.

101 In this period, Stommel’s (1961) one-dimensional laminar loop model controlled by ther-
 102 mohaline sources gave rise to the notion of switching between two different states (directions)
 103 of flow. Much subsequent investigation showed that including the wind field and more spatial
 104 dimensions greatly modified the possible inferences, but the popularity of the simple concept
 105 has been little reduced.

106 *The Modern Era and Eddying Models*

107 In the period following the Mid-Ocean Dynamics Experiment-1 (MODE), which took place in
 108 1973, it came to be recognized that the system was turbulent, on all measurable time and space
 109 scales from the full size of the ocean (10,000km) to order 1 cm, and on time-scales from seconds

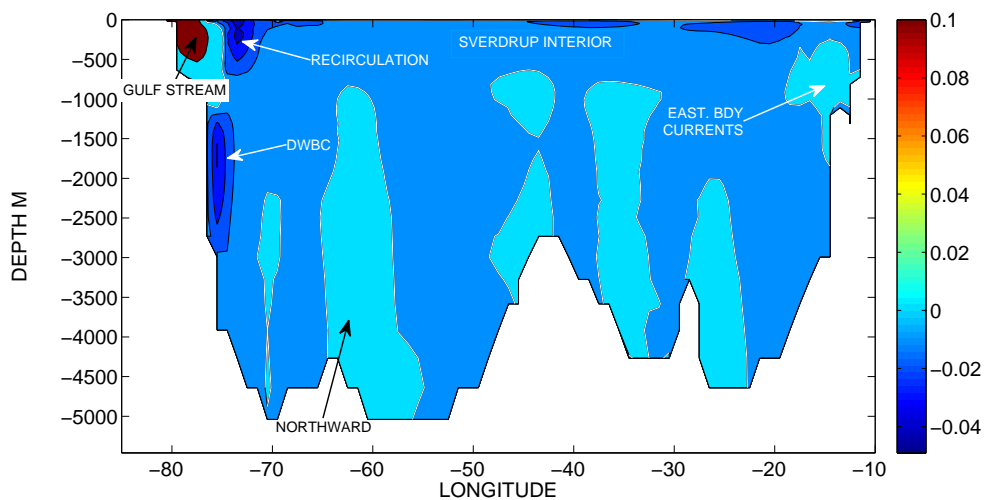


Figure 2: 19-year average meridional flow at 30°N, Wunsch and Heimbach (2013). The flow field was computed using a dynamically consistent, energy, mass, etc. conserving model, driven by known atmospheric forcing, and adjusted to be consistent with the great majority of observed data. Underlying model time step is about 1 hour. The eddy field is parameterized. The result shows the known dominant elements of the North Atlantic Ocean circulation including an intense Gulf Stream, a Deep Western Boundary Current, an interior return flow, eastern boundary currents and less well-documented interior flows over the entire water column associated in large part with the topography. Structures and volume transports vary considerably with latitude.

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110 and potentially out to the age of the fluid ocean. The parallel rise of modern computers was
111 an essential element of the modern period. MODE confirmed what many oceanographers had
112 suspected for decades—that the role of eddies and variability was much more fundamental to the
113 ocean circulation than any textbook demonstrated or explained. Accompanying the inference
114 of a fundamentally turbulent system was the recognition of the very strong, rapid, regional
115 variations in the balance of terms. In the classical period, one had subpolar and subtropical gyres,
116 equatorial region physics, eastern and western boundaries, the Southern Ocean, etc. The modern
117 era description contains all of these major regions, but with all intricately further subdivided
118 into a very long list of subregions dependent upon abyssal topographic features, recirculations
119 near eastern and western boundary currents, strongly seasonal flows, etc., rendering difficult the
120 extraction of simple overall descriptors.

121 An intuition for the complexity can be gained from viewing one of the numerous model
122 animations now available online including e.g., the one employed by the New York Times:
123 [https://www.nytimes.com/interactive/2021/03/02/climate/atlantic-ocean-climate-change.html?](https://www.nytimes.com/interactive/2021/03/02/climate/atlantic-ocean-climate-change.html?searchResultPosition=6)
124 [searchResultPosition=6](https://www.nytimes.com/interactive/2021/03/02/climate/atlantic-ocean-climate-change.html?searchResultPosition=6). Such animations are not correct in details, but do provide some intu-
125 ition as to the complexity of the real motion. The extent to which classical laminar phenomena
126 survive as major elements of the ocean circulation as it is now perceived in the modern period
127 is the subject of intense ongoing investigations, partly because they can hypothetically provide
128 an elegant description tool. Known physical elements of the variability include the spatial and
129 temporal scales of the balanced and sub-mesoscale eddies, internal waves, and likely inertial and
130 viscous sub-ranges. Energy is believed to move both towards larger and smaller scales relative
131 to the spatial scale of input (see e.g., Arbic et al., 2014).

132 A truly realistic climate model must be one of the most complicated of all pieces of non-
133 biological machinery that the human mind can encompass. A great deal has been learned from
134 model sub-components, including those that represent the ocean alone, and that model the
135 AMOC, and we are not deprecating the huge accomplishment in the modern period of global
136 and regional ocean modelling. On the other hand, a basic question that has always been present
137 concerns the realism and temporal stability of the larger laminar scales in such models when the
138 smaller scales are not resolved—and are represented by some form of simplified parameterization.

139 The ocean is a fluid, and any representation not able to include the full spectrum in space
140 and time of all motions interacting with other space and time scales must be regarded with
141 some skepticism particularly when integrated over long periods of time.. Since the time of O.
142 Reynolds, a very large literature has grown up attempting turbulence closures for very basic fluid
143 systems, often homogeneous, isotropic, unbounded. Homogeneity includes the fluid itself (e.g.,
144 no stratification) and no physics changes with position. This subject remains unfinished. The

145 ocean is inhomogeneous, anisotropic, bounded. Its most basic circulation elements include strat-
146 ification, rotation, complicated three-dimensional topography, complicated mechanical, thermal,
147 and mass forcings. To the extent that any representation of the ocean e.g. at 1 °of latitude
148 and longitude has, through parameterizations, successfully been able to quantitatively com-
149 pute large-scale quantities such as an MOC, *a quite remarkable turbulence closure* has been
150 accomplished—surely a landmark in the history of fluid dynamics!

151 **2 Discussion**

152 Curiously, invocation of AMOC shutdown, a favorite *deus ex machina* of climate science, implies
153 in the most obvious and simplest physics, an ocean circulation resembling that of the classical
154 period of physical oceanography, but inevitably encountering the complexities now known to
155 exist in the modern period ocean view. In the modern period, ocean flows are known to be fully
156 three-dimensional and turbulent, and the classical quasi-steady, laminar, upper-layer flow con-
157 trolled by a time-mean wind stress curl has become a will-o'-the-wisp—to be sought, hopefully,
158 in long time-average estimates of the ocean circulation over unknown durations.

159 An AMOC estimate must integrate across a wide-variety of features (Fig. 2) Determining
160 the amount of heat transported poleward by the circulation (the major focus of most AMOC
161 discussions, albeit usually only implicit) is a complicated matter, one in which the time required
162 to obtain a stable average is likely to vary enormously depending upon the property and the
163 latitude. In the classical view, property transports of mass, heat, etc. can increase or decrease
164 with the wind system, air-sea exchanges of enthalpy, carbon, etc. all in the “shutdown” state.

165 Time-averaging, as in Fig. 2, greatly simplifies the instantaneous flows, although the av-
166 eraging process involves an intricate product of terms in the underlying hydrodynamical/-
167 thermodynamic equations. An interesting question—one not now answerable—is whether av-
168 eraging over some long, unknown, time interval produces a classical laminar ocean circulation,
169 one that is describable from the classical time-independent equations (e.g., does Sverdrup bal-
170 ance emerge from a 100-year average?). Representing the flow and its influence by cartoons
171 carries one far from the science of oceanography.

172 Discussion of the heat carried by the ocean, or a small part of the ocean, makes little sense
173 except in the overall context of the Earth’s heat budget. Even when undergoing a global warming
174 or a cooling as into an ice age, the imbalance of incoming and outgoing radiation remains a very
175 small part of the net solar incoming flux. Modern warming rate estimates never exceed a few
176 W/m^2 at most, but the net incoming solar radiation is approximately $1400\text{W}/\text{m}^2$ (e.g., Saltzman,
177 2002). Thus, even when entering or leaving an ice age or a tropical period, the imbalance of

178 incoming and outgoing radiation remains very small compared to the overall thermal forcing.

179 If it is true that a shutdown reduces the poleward high latitude transport of energy by the
180 ocean, one can expect at zero-order that the atmosphere—globally—will tend to compensate
181 it. Bjerknes (1964) dealt specifically with the Atlantic Ocean. But the system has many ways
182 to maintain near-balance globally of incoming and outgoing radiation if the North Atlantic
183 transport of heat should shift. Changes can occur elsewhere in the oceanic poleward transport
184 (in either hemisphere), in the atmospheric transport, in the nature and degree of cloud cover,
185 surface albedo, etc. That a reduction in the AMOC heat transport would cool the northern
186 hemisphere as some like Broecker have insisted, is far from clear—the atmosphere there may
187 well warm so as to increase its radiation back to space.

188 The AMOC is one element of an immensely complicated system determining climate. Sug-
189 gesting that the entire climate system can be understood from variations in the AMOC is a bit
190 like doing an intensive study of a carburetor on an internal combustion engine, and then simply
191 assuming that all the behavior of an automobile so-powered can be inferred from its character-
192 istics. A carburetor is, per se, an extremely important part of a car, but hardly obviously the
193 determining factor in all of its complex behavior when driven.

194 Curiously, the invocation of an AMOC shutdown, as an anomalous climate state, implies that
195 the old classical understanding of the ocean circulation has become applicable. What seems to
196 be ignored, in contrast to the implications of the label, is that massive amounts of fluid are
197 moving north and south, conveying not just mass but also net amounts of heat, freshwater,
198 carbon etc.

199 *Acknowledgement.*

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